



AIR FORCE CAMBRIDGE RESEARCH LABORATORIES L. G. MANIECON PIELD, MEDFORD, MASSACHISETTS

Vertical-Attenuation Model With Eight Surface Meteorological Ranges 2 to 13 Kilometers

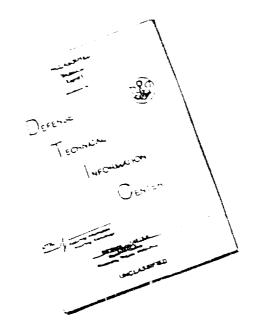
LOUIS ELTERMAN

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OPTICAL PHYSICS LABORATORY

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AIR FORCE CAMBRIDGE RESEARCH LABORATORIES

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Abstract

An examination of the haze regime shows that: (1) the aerosol properties of a surface meteorological range generally affect a mixing layer to 5 km altitude, and (2) the lower and upper visibility limits of a haze regime are defined by meteorological ranges 1.2 km and 15 km respectively. Within these limits eight meteorological ranges are selected for developing uv, visible, and ir aerosol attenuation coefficients. An aerosol scale height is derived for each meteorological range. Finally, the computed aerosol attenuation coefficients are presented as tabulations, which include previously published attenuation parameters (aerosols, molecules and ozone) to 50 km altitude.

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Symbols

 $\tau_{\mathbf{r}}'$

```
d
         Horizontal path length (km)
Hg
         Aerosol scale height (km)
         Altitude (km)
h
         Aerosol index of refraction
m
         Aerosol number density (cm<sup>-3</sup>)
Np
         Constant proportional to total number of particles between r, and r2
N_{o}
         Molecular number density (cm<sup>-3</sup>)
N_r
V_{\eta}
         Meteorological range (km)
r
         Particle radius (microns)
         Horizontal transmission
T_h
         Transmission between sea level and altitude h
T_{0-h}
         Transmission between altitude h and space
T<sub>h-∞</sub>
         Transmission between two altitudes above sea level
T_{\Delta h}
         Atmospheric ozone absorption coefficient (km<sup>-1</sup>)
\beta_3
         Aerosol attenuation coefficient (km<sup>-1</sup>)
βp
         Rayleigh (molecular) attenuation coefficient (km<sup>-1</sup>)
\beta_r
         Extinction coefficient (km<sup>-1</sup>)
^{\beta}ext
         Zenith angle
·θ
λ
         Wavelength (microns)
         Aerosol scattering cross section (cm<sup>2</sup>)
         Rayleigh scattering cross section (cm<sup>2</sup>)
\sigma_r
         Ozone optical thickness from sea level to altitude h (0-h)
	au_3
        Ozone optical thickness from altitude h to space (h-∞)
\tau_3^{\prime}
         Aerosol optical thickness from sea level to altitude h (0-h)
         Aerosol optical thickness from altitude h to space (h-∞)
\tau_{\mathbf{p}}^{\prime}
         Rayleigh optical thickness from sea level to altitude h (0-h)
	au_{	extbf{r}}
        Rayleigh optical thickness from altitude h to space (h-∞)
```

- $\tau_{\rm ext}$ Extinction optical thickness (molecular + ozone + aerosol) from sea level to altitude h (0-h)
- $\tau'_{\rm ext}$ Extinction optical thickness (molecular + ozone + aerosol) from altitude h to infinity (h- ∞)
- Ψ Aerosol size distribution function

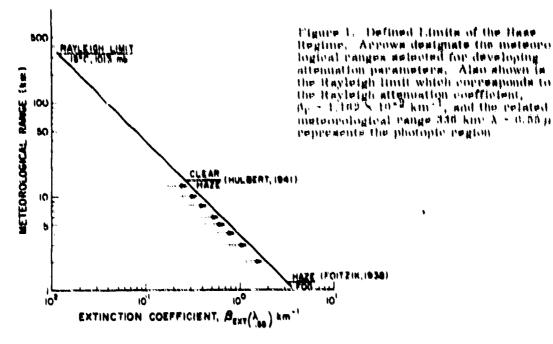
Vertical-Attenuation Model With Eight Surface Meteorological Ranges 2 to 13 Kilometers

1. INTRODUCTION

A series of atmospheric attenuation parameters which vary with wavelength and altitude are useful for carrying out a variety of exploratory calculations. Such information can take the form of curves, tabulations, or analytic expressions. It is recognized, however, that limitations exist due to variability of the atmosphere's constituents, expecially the aerosol content of the lower troposhere, which contributes extensively to the optical thickness. For example, in the photopic region, assuming a representative wavelength λ = 0.55 μ and a meteorological range of 23 km near the surface, the aerosol content in the first 3 km above sea level accounts for about 70 percent of the total optical thickness. If surface conditions are hazy or polluted, the aerosol content accounts for a larger percentage. This suggests that the treatment of atmospheric attenuation can be improved by introducing aerosol parameters related to the easily measured meteorological range, that is, by introducing quantitatively a haze regime which is considered as encompassing a series of meteorological ranges between those associated with normally clear conditions and fog. See Figure 1.

Near the surface and at low altitudes, the aerosol constituent is ubiquitous and highly variable, and an aerosol component is present even for a very clear atmo-

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spheric condition. Thus, the boundaries of the haze regime (used in the sense of diminished meteorological range) must be defined. The haze limits will be based on the Roschmieder (1924) definition,

$$V_{\eta} = \frac{3.91}{\beta_{\rm ext}} , \qquad (1a)$$

, and the related

where

$$\beta_{\text{ext}} = \beta_{\text{p}} + \beta_{\text{r}}$$
 (1b)

 V_{η} is the meteorological range (km) and β_{ext} , β_{r} , β_{p} are the extinction, Rayleigh, and aerosol attenuation coefficients (km⁻¹) respectively. Equations (1a) and (1b) apply to sea level conditions and the photopic region represented by $\lambda = 0.55 \,\mu$. When converted to a log-log trace (Figure 1), several boundaries can be designated conveniently. The Rayleigh limit, 336 km, is based on the standard atmosphere (15°C, 1013 mb). Foitzik (1938) found that the haze-fog transition is relatively abrupt, and, therefore, can be readily identified; and that $V_n \approx 1.2 \text{ km}$ represents the transition. Folisik's result is confirmed adequately by Nethurger and Chien (1960), Hullrich (1963), Closs (1963, 1964) and the analysis by Eldridge (1969).

In contrast, the literature pertaining to a boundary condition between a clear and hasy atmosphere is meager. Since no abrupt changes occur in this region, the requirement is tantamount to dealing quantitatively with psycho-physical observations. Hulbert (1941) correlated meteorological range with atmospheric conditions such as dense fog, light fog, hase, clear, and so forth, using a telescopic photometer, the measurements being made in the vicinity of Washington, D. C. His results led him to propose $V_{\eta} = 15$ km as the base-clear boundary condition. Despite the subsective element in this result, it provides some guidance. In conjunction with Foitsik's observations, it permits defining the base regime as 1.2 $\lesssim V_{\eta} \lesssim 15$ km. The meteorological ranges and corresponding parameters shown in Figure 1 and Table 1 will be used in the material to follow because they are spaced at convenient

Table 1. Meteorological Ranges and Corresponding Parameters (Representative Photopic Wavelength $\lambda = 0.55 \mu$)

V _ŋ (km)	$\frac{\beta_{\mathbf{e} \times \mathbf{t}}}{(\mathbf{k}\mathbf{m}^{-1})}$	β _r (km ⁻¹)	$\frac{\beta_p}{(km^{-1})}$	H _p (km)
 2	1,955	0,0116	1,043	0,84
3	1,303	0.0116	1,291	0,90
4	0,978	0.0116	0,966	0,95
5	0,782	0,0116	0.770	0,99
6	0,652	0,0116	0,640	1.03
8	0,489	0,0116	0,476	1.10
10	0,391	0,0116	0,379	1,15
13	0.301	0,0116	0,289	1.23
ν _η - m	ateorological ra	unge		
•	tinction coeffici	ent	•	
	ayleigh attenuati	on coefficient		
β _p - ae	rosol attenuatio	n coefficient		
· ·	rosol scale heig	ght		

logarithmic intervals, are adquately separated from the haze-fog transition, and are within the haze regime characterized by diminished meteorological range. The scale height in the last column of the tabulation will be discussed later.

2. SPECTRAL METEOROLOGICAL RANGES

The concept of photopic meteorological range can be widened spectrally if concurrent measurements of attenuation coefficients at other wavelengths are available. The work of Curcio, Knestrick, and Cosden (1981), which is the basis for Figure 2, is an example where $\beta_p(V_4,\lambda)$ was obtained through a series of concurrent measurements. Because of the quantity of data obtained, the results for the meteorological range $V_{\eta}=4$ km are considered representative by these authors. A family of distributions, $\beta_p(V_{\eta},\lambda)$, can be computed if the $\beta_p(V_4,\lambda)$ values are used in conjunction with Eq. (1) so that

$$\beta_{p}(V_{\eta}, \lambda) = \beta_{p}(V_{4}, \lambda) \cdot \left[\frac{3.91}{V_{\eta}} - \beta_{r}(\lambda_{.55})\right] / \left[\frac{3.91}{V_{4}} - \beta_{r}(\lambda_{.55})\right] ,$$
 (2)

 V_{η} (km) being the photopic (λ = 0.55 μ) meteorological ranges of interest. Using Eq. (2), the aerosol attenuation coefficient is found for various combinations of meteorological range and wavelength, 0.27 to 2.17 μ (Table 2), that is, 160 surface values, $\beta_{p}(V_{\eta}, \lambda)$. The shapes of the distributions so determined (Figure 2) conform rigorously to the distribution for V_{η} = 4 km, on which they are based.

Because of the functional importance of Eq. (2), it would be in order to examine its implications, especially those related to particle size considerations. If we

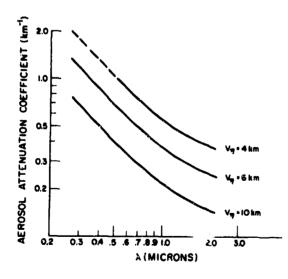


Figure 2. Wavelength Distributions of the Aerosol Attenuation Coefficient for V_η = 6 km and V_η = 10 km Derived From V_η = 4 km Using Eq. (2). The V_η = 4 km curve is obtained from measurements by Curcio, Knestrick and Cosden (1961), which included the wavelength region $0.40 \le \lambda \le 2.17\mu$. An extrapolation to 0.27μ permits computations for an overall 20 selected wavelengths, $0.27 < \lambda < 2.17\mu$, and 8 meteorological ranges $2 \le V_\eta \le 13$ km. The dash portion of the top curve represents extrapolation

consider a real atmosphere the aerosol sizes within unit volume determine the aerosol attenuation coefficient described by

Table 2. Surface Aerosol Attenuation Coefficients Corresponding to Figure 2

		$\beta_{p}(h_{o}, \lambda, V_{\eta})$	
λ (μ)	V ₄	V ₆	V ₁₀
0,27	2,00	1.33	7.85×10 ⁻¹
0.28	1.89	1.25	7.42×10 ⁻¹
0,30	1.78	1.18	6.98×10 ⁻¹
0.32	1.67	1.11	6.55X10 ⁻¹
0.34	1.56	1.03	6.12×10 ⁻¹
0.36	1.45	9.61×10 ⁻¹	5,69X10 ⁻¹
0.38	1.40	9.28×10 ⁻¹	5.49×10 ⁻¹
0.40	1.30	8.61X10 ⁻¹	5.10×10 ⁻¹
0.45	1.15	7.62×10 ⁻¹	4.51×10 ⁻¹
0.50	1.05	6.96×10 ⁻¹	4,12×10 ⁻¹
0.55	9.66×10 ⁻¹	6.40×10 ⁻¹	3.79×10 ⁻¹
0.60	8.60×10 ⁻¹	5.70×10 ⁻¹	3.37×10 ⁻¹
0.65	7.80×10 ⁻¹	5.17×10 ⁻¹	3.06X10 ⁻¹
0.70	7.30×10 ⁻¹	4.84×10 ⁻¹	2.86×10 ⁻¹
0.80	6.40×10 ⁻¹	4.24×10 ⁻¹	2.51×10 ⁻¹
0.90	5.80×10 ⁻¹	3.84×10 ⁻¹	2.28×10 ⁻¹
1.06	5.20×10 ⁻¹	3.45×10 ⁻¹	2.04×10 ⁻¹
1.26	4.70×10 ⁻¹	3.11×10 ⁻¹	1.84×10 ⁻¹
1.67	4.00×10 ⁻¹	2.65×10 ⁻¹	1.57×10 ⁻¹
2.17	3.60×10 ⁻¹	2.39×10 ⁻¹	1.41×10 ⁻¹

$$\beta_{p}(m, \lambda) = \int_{r_{1}}^{r_{2}} \sigma_{p}(m, r, \lambda) n(r) dr, \qquad (3)$$

$$n(r) = N_0(V_n) \psi(r)$$
 (4)

and when combined

$$\beta_{p}(\mathbf{r},\lambda,\mathbf{V}_{\eta}) = N_{o}(\mathbf{V}_{\eta}) \int_{\mathbf{r}_{1}}^{\mathbf{r}_{2}} \sigma_{p}(\mathbf{r},\lambda) \psi(\mathbf{r}) d\mathbf{r}. \qquad (5)$$

In these expressions, β_p is the aerosol attenuation coefficient; the index of refraction is m (to be omitted following Eq. (3) because subsequent considerations will assume m invariable); \mathbf{r}_1 and \mathbf{r}_2 are the lower and upper radii limits of the size distribution $\mathbf{n}(\mathbf{r})$; \mathbf{N}_0 is a constant proportional to the total number of particles between \mathbf{r}_1 and \mathbf{r}_2 ; ψ is the size distribution function (the same for all selected meteorological ranges).

It is implicit in Eq. (5) that β_p and correspondingly the aerosol number density determine the meteorological range (V_{η}) . The integral in Eq. (5) is a wavelength function independent of the meteorological range, which accounts for the conformity in shape of the curves in Figure 2.

3. STATEMENT OF OBJECTIVES

The material thus far has dealt with: (1) the limits of the haze regime (in terms of meteorological range), (2) derivation of spectral aerosol attenuation coefficients $(0.27 \le \lambda \le 2.17\mu)$ for a series of meteorological ranges, and (3) an examination of the assumptions implicit in the derivation of these coefficients. Now, a statement of objectives can be made. Specifically, aerosol scale heights will be determined for the coefficients in accordance with their meteorological range and their vertical distribution. Then values of the coefficients for km intervals (0-5 km) will be computed. To the coefficients will be added previously published Rayleigh, ozone, and aerosol parameters for altitudes to 50 km (Elterman, 1968) in order to formulate an attenuation model for a haze regime with eight meteorological ranges $(2 \le V_{\eta} \le 13 \text{ km})$.

4. AEROSOL MIXING LAYER

The procedure for assessing the aerosol attenuation coefficient above the surface can parallel that used for a clear atmosphere (Elterman, 1968), which entails the ap-

plication of a suitable aerosol scale height. As an introduction to scale height considerations, it is noted that, meteorologically, a significant role is assigned to the altitude interval up to several km above the surface. This is a region of strong vertical mixing determined by such factors as heat-transfer across the earth-air interface, advective winds, and consequent turbulence attributable to the region's topography. The resultant vertical flow, mechanical and convective, is characterized, meteorologically, as a mixing depth equivalent to the vertical extent of the mixing layer. When dealing with aerosol attenuation coefficients, aerosol conditions in this layer can be examined in terms of mixing ratios such that for a selected altitude h

$$\frac{\beta_{\mathbf{p}}(\mathbf{h},\lambda)}{\beta_{\mathbf{r}}(\mathbf{h},\lambda)} = \frac{\sigma_{\mathbf{p}}(\lambda)}{\sigma_{\mathbf{r}}(\lambda)} \cdot \frac{N_{\mathbf{p}}(\mathbf{h})}{N_{\mathbf{r}}(\mathbf{h})}, \qquad (6)$$

where β_p and β_r are respectively the aerosol and Rayleigh attenuation coefficients (cm $^{-1}$); and N_p and N_r are respectively the aerosol and molecular number densities (cm $^{-3}$). The terms σ_p and σ_r , which are respectively the aerosol and Rayleigh cross sections (cm 2), tend to remain constant with altitude (a reasonable assumption). Eq. (6) then asserts that $\beta_p(h,\lambda)$ / $\beta_r(h,\lambda)$, known as the optical mixing ratio, is proportional to the number density mixing ratio $N_p(h)$ / $N_r(h)$. The size distribution for N_p comprises aerosols sufficiently small to be responsive to the usual factors conducive to mixing. Meteorologically, the mixing depth is considered to be 3 km or less, so that an aerosol mixing depth determined from an optical mixing ratio or a number density mixing ratio or even by direct (in-situ) measurement of $N_p(h)$ should be in agreement.

The conclusion, however, based on aerosol measurements sufficient to provide a meaningful average, is that the mixing depth normally extends to a greater altitude. Siedentopf's (1944) sky luminance measurements (18 aircraft flights) show that on the average, the aerosol concentration decreases exponentially with altitude and that the scale height undergoes a significant change between 5 and 6 km. Penndorf's (1954) analysis of solar attenuation observations (8 aircraft flights) shows the scale height change to occur at 4.5 km. An examination of Rosen's (1967) balloon photoelectric countermeasurements, selecting only those profiles where the surface layer is readily discerned (37 profiles obtained on ascent, descent and for 2 wavelengths), shows that the average mixing depth is 5.4 km. An analysis of optical probing measurements (Elterman, Wexler, and Chang, 1969) yielded optical mixing ratios which show that the depth of the surface layer averages 5.3 km (79 profiles at 0.55 μ wavelength). Blifford and Renger (1969) completed a series of 22 aerosol collections using an aircraft-mounted impactor. Samples obtained to 9.1 km provide mixing ratios that indicate the mixing depth to be in the range 3-6 km.

Measurement of the atmospheric aerosol distribution has received considerable emphasis in the USSR, for example, in the work of Faraponova (1965), who conducted more than 60 aircraft flights in cloudless weather (solar atmospheric attenuation) to 6.5 km altitude, and in the summary by Kondratiev (1969). In general, the USSR findings are compatible with those previously described.

An overall assessment of the results shows that the aerosol content for the low altitudes is characterized by a mixing depth between 4.5 and 5.5 km. As has been mentioned, it is somewhat higher than the mixing depth of the meteorology discipline. However, the difference is understandable when it is considered that in almost all instances, aerosol measurements were conducted over the continent and with cloudless skies, whereas the mixing layer in the meteorological sense represents less limited conditions. Accordingly, the designation "aerosol mixing layer" will be used, and will be assigned a depth of 5 km (considered representative). Within the aerosol mixing layer, considerable variation and stratification (frequently due to inversions) can occur but, on the average, the particle distribution, $N_{\rm p}(h)$, and correspondingly the aerosol attenuation coefficient, $\beta_{\rm p}(h)$, decreases exponentially for the altitude region 0-5 km. The rate of decrease can be expressed as a constant scale height although, as will be shown, not necessarily the same scale height for each meteorological range. At 5 km, effects of mixing are substantially diminished.

5. THE USE AND JUSTIFICATION FOR A SCALE HEIGHT FAMILY

If aerosol conditions at the upper terminus of the aerosol mixing layer are relatively stable compared to those at lower altitudes, as discussed, it suggests that the scale height characterizing the aerosol mixing layer is related to the meteorological range. The existence of such a relationship was examined by means of Figure 3, using $\lambda = 0.55\mu$. Specifically, the scale heights were determined by: (1) utilizing the surface values, β_p ($\lambda_{.55}$, V_η), for the meteorological ranges of interest derived from Eq. (2); and (2) taking from published tabulations (Elterman, 1968) at 5 km the aerosol attenuation coefficient, β_p (h_5 , $\lambda_{.55}$) = 5.0 × 10⁻³ km⁻¹. This quantity (assumed representative because the tabulations are based on 79 sets of measurements) is considered relatively independent of the meteorological range for reasons already given. With surface values and the 5 km value established, the aerosol scale height (H_p) was derived for each meteorological range by using

$$\beta_{p} (h_{5}, \lambda_{.55}) = \beta_{p} (h_{o}, \lambda_{.55}) e^{-h/H} p$$
(7)

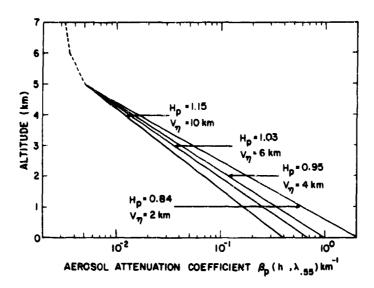


Figure 3. Relationships of Four Aerosol Scale Heights (H_p) With Meteorological Ranges (V_n), Aerosol Attenuation Coefficients (β_p) and Aerosol Mixing Layer Altitudes (0-5 km). The aerosol scale height family was computed using λ = 0.55 μ . The dash line represents values of β_p (h) above 5 km (Elterman, 1968)

It is implicit in this procedure that the derived scale height family is independent of wavelength. With H $_p$ known, the values for $\beta_p(h,\lambda)$ can be calculated (using Eq. (7)) for each km interval up to 5 km (the aerosol mixing layer). To summarize, the decrease of β_p with altitude is represented by a family of scale heights and each scale height depends on the meteorological range of interest (Table 1).

The validity of this procedure depends not only on the value but also on the variability of $\beta_p(h_5,\lambda)$ as it affects the related parameters of interest. In this respect, we note that the aerosol optical thickness up to 5 km, $\tau_p(h_{0-5},\lambda)$, is an important objective; and also that it is obtained by integration, which makes it sensitive to change of its composite elements, especially at low altitudes. Thus, a suitable evaluation of the aerosol scale height family derived in accordance with Figure 3, is to vary $\beta_p(h_5,\lambda_{..55})$ significantly and to examine the optical thickness and transmission change for the first 5 km and for the several meteorological ranges indicated. This was done by changing $\beta_p(h_5,\lambda_{..55})$ = 5.0 × 10⁻³km⁻¹ by a standard deviation, σ = ± 3.4 × 10⁻³. For this calculation, as previously mentioned, the mean β_p and σ values at 5 km altitude were obtained from 79 selected optical probing measurements (Elterman, 1968). The resulting aerosol optical thickness and transmission changes are shown in Figure 4. For the eight meteorological

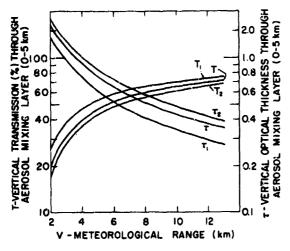


Figure 4. Comparison of Vertical Optical Thickness and Transmission for $\pm\,\sigma$ (Standard Deviation) and a Mean Aerosol Attenuation Coefficient $\beta_{p}\,(h_{5}\,,\lambda_{.55})$ at the Top of an Aerosol Mixing Layer Having 5 km Depth. T and τ represent the transmission and optical thickness based on $\beta_{p}\,(h_{5}\,,\lambda_{.55})$ = $5.0\times10^{-3} {\rm km}^{-1}$, the mean of 79 measurements. T_{1} and τ_{1} represent the transmission and optical thickness based on $\beta_{p}\,(h_{5}\,,\lambda_{.55})$ - σ , where σ = 3.4×10^{-3} . T_{2} and τ_{2} represent the transmission and optical thickness based on $\beta_{p}\,(h_{5}\,,\lambda_{.55})$ + σ

ranges, a $+\sigma$ of the aerosol attenuation coefficient corresponds to an average of 3.6 percent change in vertical transmission (attributable to aerosols only). Similarly, a $-\sigma$ corresponds to a 6.4 percent vertical transmission change. The changes are unequal due to exponential relationships. These changes are relatively modest especially when it is considered that the aerosol attenuation coefficients representing the meteorological ranges 2 km to 13 km extend over an order of magnitude. Based on this discussion, it is concluded that for conditions of diminished meteorological range in the aerosol mixing layer, there is sufficient justification for the use of an aerosol scale height family, and that the scale height selected is best determined from the meteorological range of interest.

6. THE OPTICAL TRICKNESS COMPUTATION

Since the aerosol attenuation coefficients, β_p (h, λ , V_η), as well as the aerosol scale heights, H_p (constant for each meteorological range) are known, an analytical expression for the aerosol optical thickness, τ_p , was derived as follows:

$$\tau_{p}(h,\lambda,V_{\eta}) = \int_{0}^{h} \beta_{p}(h,\lambda,V_{\eta}) dh.$$
 (8)

For a given wavelength, the aerosol scale height expression applicable to the haze regime depends on the meteorological range and the vertical distribution of the aerosol attenuation coefficient. Accordingly,

$$\beta_{p}(h,\lambda,V_{\eta}) = \beta_{p}(h_{o},\lambda,V_{\eta}) e^{-h/H_{p}(V_{\eta})}.$$
(9)

Combining Eqs. (8) and (9) and integrating,

$$\tau_{\mathrm{p}}\left(\mathsf{h},\lambda,\mathsf{V}_{\eta}\right) = \mathsf{H}_{\mathrm{p}}\left(\mathsf{V}_{\eta}\right) \cdot \beta_{\mathrm{p}}\left(\mathsf{h}_{\mathrm{o}},\lambda,\mathsf{V}_{\eta}\right) - \mathsf{H}_{\mathrm{p}}\left(\mathsf{V}_{\eta}\right) \left[\beta_{\mathrm{p}}\left(\mathsf{h}_{\mathrm{o}},\lambda,\mathsf{V}_{\eta}\right) \cdot \mathrm{e}^{-\mathsf{h}/\mathsf{H}_{\mathrm{p}}\left(\mathsf{V}_{\eta}\right)}\right]. \tag{10}$$

Applying Eq. (9) to the bracketed factor in Eq. (10),

$$\tau_{p}(h,\lambda,V_{\eta}) = H_{p}(V_{\eta}) \left[\beta_{p}(h_{o},\lambda,V_{\eta}) - \beta_{p}(h,\lambda,V_{\eta})\right]. \tag{11}$$

Equation (11) was used to compute the aerosol optical thickness for the combinations of wavelength, altitude, and meteorological range required for Figure 4, and the twenty tabulations (surface to 5 km) in Tables 3.1 to 3.20.

7. SUMMARY AND CONCLUDING REMARKS

To formulate an atmospheric attenuation model with meteorological ranges for a haze regime, it was necessary first to define the limits of the haze regime. Following this, eight meteorological ranges were selected within the regime and the surface aerosol attenuation coefficient distribution in wavelength was determined for each meteorological range. From the surface aerosol attenuation coefficients, the $\beta_p(h,\lambda,V_\eta)$ then were computed for km intervals to 5 km altitude (typical depth of the aerosol mixing layer) by applying a scale height derived for each meteorological range. Optical thickness values required for the model were computed with Eq. (11). Finally, parameters from an earlier published attenuation model (Elterman, 1968) were combined with those derived in this paper, in order to provide continuity to 50 km altitude.

The shortest wavelength used in this model is 0.27 microns. The use of shorter wavelengths would required the treatment of $\rm O_2$ absorption and its attendant uncertainties. The longest wavelength used is 2.17 microns. Calculations for longer wavelengths are complicated by the presence of absorption bands of $\rm H_2O$, $\rm CO_2$ and their wings. Also, at longer wavelengths and low altitude haze conditions, absorption by the aerosol itself is an unknown factor. In between, a total of 20 wavelengths is chosen, within the atmospheric windows and for the ultraviolet region where ozone absorption is important (Table 2). If required, a satisfactory interpolation for the optical parameters can be made between wavelengths in the region 0.27 to

about 1μ because light extinction in this spectral region is caused primarily by scattering and ozone absorption and both processes are slowly varying functions of the wavelength. This is true of the extinction coefficients, $\beta_{\rm ext}(\lambda)$, as well as their components, $\beta_{\rm p}(\lambda)$, $\beta_{\rm r}(\lambda)$, and generally $\beta_{\rm 3}(\lambda)$.

Beyond 1μ , the computations did not include molecular absorption. Therefore, the tables for wavelengths 1.06, 1.26, 1.67, and 2.17 μ represent the IR windows only. The presence of absorption bands due to $\rm H_2O$ and other gases does not permit interpolation between $1.06 \le \lambda \le 2.17\mu$, the near IR region considered in this model, unless the interpolation is limited to the Rayleigh and aerosol parameters (no ozone absorption present).

8. TABULATIONS

The tabulations that follow are in computer notation. For example, read $5.96 - 2 = 5.96 \times 10^{-2}$ and $5.96 \cdot 2 = 5.96 \times 10^{2}$.

The format deals systematically with a multiplicity of variables, thus permitting a variety of exploratory calculations. As an example, the extinction coefficients can be used for exploratory transmission calculations. The atmospheric extinction coefficient is the sum of all the attenuating components:

$$\beta_{\text{ext}}(h,\lambda,V_n) = \beta_{\text{r}}(h,\lambda) + \beta_{3}(h,\lambda) + \beta_{p}(h,\lambda,V_n).$$
 (12)

For horizontal transmission over a path length (d) at selected altitude, wavelength, and meteorological range

$$T_{h}(h,\lambda,V_{\eta}) = \exp \left[-\beta_{ext}(h,\lambda,V_{\eta}) \cdot d\right]. \tag{13}$$

For vertical and slant path transmission from sea level to an altitude of interest at zenith angle θ ,

$$T_{o-h}(h,\lambda,V_{\eta}) = \exp\left[-\tau_{ext}(h,\lambda,V_{\eta}) \cdot \sec\theta\right].$$
 (14)

For vertical and slant path transmission between two altitudes (\mathbf{h}_1 and \mathbf{h}_2) above sea level,

$$T_{\Delta h}(h,\lambda,V_{\eta}) = \exp -[\tau_{\text{ext}}(h_2,\lambda,V_{\eta}) - \tau_{\text{ext}}(h_1,\lambda,V_{\eta})] \sec \theta. \tag{15}$$

For a vertical and slant path transmission from a selected altitude out into space,

$$T_{h-\infty}(h,\lambda,V_{\eta}) = \exp\left[-\tau_{\text{ext}}^{\dagger}(h,\lambda,V_{\eta}) \sec \theta\right].$$
 (16)

When used individually, Rayleigh, aerosol, and ozone parameters are formulated similarly.

				• '	Table 3. i.	Param	Parameters at	0.27 Microns	rons				
j (Alt.	Rayleigt.	Rayleigh	Layleigh	Aerosol	Aerosol		Ozone	22000	i de	3		
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	n →	2.242 -1	90.	1.928	0 70-	8	3.342	7-66 -1	. 390	70.434		i	16.27
	^		**	1.513		25-378	1.0	1- 50-6	.714	74. X			73.215
	•		. 563	1.335				1- 61-9	1.566	本	1-16 .	£;	71.478
(•		.753	1.176		378		7- 67-6	1.136	49-65		. y	Z. Z.
7	<u>ب</u>		***	1-040		104.4			664.7			5.5	3.5
l	3 :		1.422	ž.		3.333		7		Ç.		•	2.7
			1.696	167-		3.36	-032	2.81	12.361				
			779-1	8		3.302	•10-	3.5	25-765	3			
			710			3.363	-683	3.76	65.181	Z.X.			
	3	1-913 -4	1.926	200		3-342	9	1- 50-6	1122	1.576		-	,
				į		3.54	2	3-91 -2	12.131	• 828		. 4	128-
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	`		*1**	1.513			Ţ.	-	.716	73-240		2.5	72,742
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•	•		.753	1-174		7.327			976	2.63		1	7.25
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,			1.422	765		9, 1	ī,	7- 11-	7-985	156-13		*	12.5
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				3		7	3	3.81 -2	73.13;	Š			i.
	7.	2-747	36	1.926		86	1.67			78 87	7	ļ	,
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		1-673-1	*14.	115-1		1.350	-279		4.6.7	44.4			
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u	, ,	1- 7/4-7		***		1.522	.197		929-7	**.52	7.7	121	
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817	775	. 000 . 2117 . 414 . 753 . 753 . 1642	1.822	.000 .217 .414 .593 .753 .898 .1992 .1992 .1997 .1917	.000 .217 .214 .513 .753 .846 .1.472 .1.472 .1.472 .1.472 .1.473
	1011	2.725		2.7.2 2.03.1 1.7.3.2 1.7.3.2 1.7.3.3 1	100 100 100 100 100 100 100 100 100 100
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			ŀ		Table 3.	2. Param	neters a	t 0.28 Microns	crons				
Het.	Alt.	Rayleigh	Rayleigh	Rayleigh	Aerosol	Aerosol	Aerosol	Ozone	Одове	Czone	Er.	H	5
		coeff	thick.	thick.	coeff	rhick	thick	. Coeff	Trick.	100		1011	
E	(<u>k</u>	(k=-1)	(q-0)	Ť	(km ⁻ 1)	(0-h)	Ė	- I	(q-p)	() ()	 	(10)	(
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	2	1- 896-1	000	1.043		000	1-197		90(35.416		٤	3
	-	1-167-1	991.	1.460		761-7	1-000	_	355	35.455		2.7	37.914
	٧.	1-000	3.5	1-2-1		2-847	.350		2.590	35.126		3.890	36.768
	- .	- 944-1	.506	1-139		3-039	.156	_	. 317	34.839		4.522	36-136
6	* r	1. 303 -1	197	1.002	2.90 -2	3.055	101	1- 04-2	062-1	34.586	3.99 -1	1.86	35.690
1		6.374 -2	1.214			3.122	450		7.1.	34-350			37.313
	ć.	4-096 -2	1.448	177		3.167	•029	-	245.5	29.274		15.157	79.501
	3	1-414-5	1.555	0.00		3.187	010	_	13.305	22.911		17.741	22.911
	ζ:	6.373 -3	1.603	.042		3.194	•003		23.312	12.004		27.809	12.849
	ç 7	4.146 -5	1.636	60.		3-156	000		33.506	2-310		38.338	2.329
	2	t- 150-1	***			3-157	000		35.306	010.			.
	σ.	1-946-1	000	1.645		0007	2.299		. 300	35.816	3.13 0	3	39.760
	- -	1- /9/-1	991	1.460		.208	.790		.361	35.455	_	2.056	37.704
		1-000-1	* 5	1.231		1.994	306		JE 4.	35-126		3.036	36.724
		1-303 -1	244	1 000		951.6	2		206.	34.839		3.631	36.128
(~)		1-171 -1	747	879		2.214	1010		062-1	94.75		1.0.4	35.689
•	9	6.314 -2	1.214	164.		2.244	-055		2.74	33-062	_		33.528
	:	3-0-6	1.448	141.		2.269	.029		6.542	29.274	_	10.256	29.501
	? .	2- 414-1	1.555	060		5.289	020-		13.305	22.811	-	16.869	22.711
	Ç	6- 575 -5	1.603	7,0		5.2%	•003		23.312	15.804		26.931	12.849
	; ;	1-340-1	1.030	60	2-62-5	2.298	000	1-16-4	33.506	2.310	4.53 -1	37.440	2.320
									33.30	2	_	33.143	10.
	2	1-948 -1	000	1.645		009.	1.828		.300	35.815	_	8	39.289
	- , ·		981	1.653		1.156	.672		196.	35.455		1.703	37.585
			* 3	1.231		1.549	-279		065.	35.126	_	2.592	36.697
•	• •		900	1.002		1.053	•		776.	36.833		3.166	36-124
4	•		.167	.879		¥2.1	•082		1-167	34,353	_	1.977	35.313
	<u>.</u>		1.71	.431		1.774	-055		2.774	33.042		5.761	33.528
	2 2		844-1	161.			-029		5.542	29.67		9.789	29.501
			1.603	240		27871	000		13.705	119.77		16.378	116.22
	3.3	1-340 -3	1.636	. 000	5.62 ->	1.828	000	4.57 -1	13.206	2.310	4.58 -1	36.970	2.320
	Ž		1.644	100		1.828	000.		35.306	.010		39.278	170.
	7	1- 846-1	.000	1.665		000	1.534		7360	35 16	68.0	Ş	400
		1-707-1	.136	1.463		-636			1991	25-455	_	4	27 609
	`		**	1.2.1		1.213	1920		0 ¢ ¢ •	35,126		7.316	36.679
	-		• 500	1.139		1.392	-145		116.	34.839		2.875	36-120
M			.643	1.002		1.434	.100		1.230	34.586		3.307	35.688
•	٠:		191	6.20		655-1	• 082		1.667	34.350		3.682	35-313
	2 -		1.264	4 4 4		6/4-1	• 055		2.774	33.042		5.467	33.528
	: ?	\- +!+·!	1.355	040	2.55 -3	1.524	670-	1.74	13,105	22.811	8.4	464	29.501
	,		1.003	.042	_	1.031	.003		23.012	12.804		26.146	12.849
	Ç.	1. 146 -3	1.036	٠٥٥٠	_	1.534	.000		33.506	2.310		36.675	2.320
	č	1.633 -4	1.044	109.		1.534	000-		35.306	-010		38.984	170-

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1, 2, 2, 94 1, 150 1, 15	25.5.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.2.		. 450 . 167 . 051 . 051 . 007 . 001		6.463	960.		1111	3.295		3.520	4.137
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1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1,	1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1		. 031 . 007 . 001 . 032 . 1. 084		3.014	600		1,739	2.169		6.693	2.964
	1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1		.001 .001 1.222 1.084		3.620	.003		2.193	1.220		6	7.64
1,000	11.11.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.	,	1.222		3.023	000		3.193	.220		7.430	1227
1110 1.00	1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.1.		1.222		3.023	80.		3.112	100-		7.655	-002
	1.1121. 1.1121		1.064		000		2, 60	60.6		;		
1,000	1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00				1.424		3- 52- 6	960	24.4	2.56	8	6.809
1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1,	12 7.244 13 7.244 14 7.44 15 7.44 16 7.44 17 7.44 1		. 454		1.882			990	3.367			5.213
1.00	13 7.24 10 10 10 10 10 10 10 10 10 10 10 10 10		*		2.030			.393	3,320		117.7	97.0
1,	2.2 2.2 2.2 2.2 2.2 2.2 2.2 2.2 2.2 2.2				2.078			111.	3.295		2.672	7.136
1.	7.299 1.050 1.77 9.940 1.180 1		320		2.122			0+1-	3.273		2.802	100
1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1, 1,	4. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1.	_	.167		2-146			*564	3.148		3.288	3.521
1.00	1.073 1.073 1.073 4.693 4.693		-067		2.165			1.729	2.173		3.844	2.964
1,175	1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1. 1		160.		2.172			2-193	1.220		4.559	2.250
1.444	1.012 1.012 1.013 9.073 4.693		8		2-174			3-193	.220		6.581	1.224
1.500 1.222 1.78 0 0.000 1.730 3.59 2 0.000 3.413 1.99 0 0.000 0.251 0.000 0.252 0.000 0.0	1.444 1.312 1.158 1.073 4.672 4.681		į		611.5			3.412	100.		6.807	-005
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	1.168 1.073 9.677 6.693		1.044		160-1			000	3.413		000	6.364
1.00	724.4 724.4 724.4	_	.959		1.463			490	3.378		1.263	101.5
1	4.643	_ •	94.		1.591			. 193	3.320		2.050	4.26
10 4.081 -2 -2 -2 -2 -2 -2 -2 -	4.482	_	79		1.634			1111	3.295		2.229	4.136
1, 2, 3, 40 - 2 1, 1075 1, 10		_	.320		1.678			091-	3.273		2.358	4.007
1.00 2.00 2.00 2.00 1.00	2-749	_	.147		1.702			* 604	3.148		2.643	3.521
1.71? -4 1.221 .001 9.91 -7 1.730 .003 1.82 -1 2.193 1.220 1.67 -1 5.111	0.0.1		28.		1.72)			1.239	2.173		3-400	2.964
1.712 - 4 1.721	9.5		700		1.727			2-193	1.220		5.111	1.256
1-444 -1 -100 1.222 1-42 0 -100 1-452 3-60 -2 -300 3-413 1-60 0 -300 1.112 1 -138 1.084 5-47 -1 -886 5-56 3-29 2 -334 3-378 6-71 1 1-058 1.112 1 -138 1.084 5-47 -1 1-233 2-49 2 -316 3-347 3-39 -1 1-531 1.075 -1 -376 -846 -47 -2 1-318 -136 3-32 -2 -137 3-39 1-97 1 -184 1.075 -1 -279 -2 -279 -2 -279 -279 -279 -279 -279 1.0 -2 -2 -279 -2 -279 -279 -279 -279 -279 1.0 -2 -2 -279 -2 -279 -279 -279 -279 1.0 -2 -2 -279 -279 -279 -279 -279 -279 1.0 -2 -2 -279 -279 -279 -279 -279 1.0 -2 -2 -279 -279 -279 -279 -279 1.0 -2 -279 -2 -279 -279 -279 -279 1.0 -2 -279 -279 -279 -279 -279 -279 1.0 -2 -279 -279 -279 -279 -279 -279 1.0 -2 -279 -279 -279 -270 -270 -270 1.0 -279 -279 -279 -279 -270 -270 1.0 -279 -279 -279 -279 -270 -270 1.0 -279 -279 -279 -270 -270 -270 1.0 -279 -279 -279 -270 -270 -270 1.0 -270 -270 -270 -270 -270 1.0 -270 -270 -270 -270 -270 1.0 -270 -270 -270 -270 -270 1.0 -270 -270 -270 -270 -270 1.0 -270 -270 -270 -270 -270 1.0 -270 -270 -270 -270 1.0 -270 -270 -270 -270 1.0 -270 -270 -270 -270 1.0 -270 -270 -270 -270 1.0 -270 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.0 -270 -270 -270 1.	1.717		8		1.730			3-193	-220		6.137	.227
1-444 - .000 1.222 1-42 0 .000 1.452 3-40 2 .000 3-413 1-60 0 .000 1-114 - .158 1.084 5-07 1 .886 5-66 3-29 2 .034 3-318 6-71 1 .058 1-114 - .258 .247 2 .243 2-46 2 .256 3-347 3-39 1 .1531 1-1075 - .376 .846 6-47 2 1-318 2-56 2-66 3-347 3-39 1 .1531 1-1081 - .487 .487 .243 .243 .243 .243 .243 .243 .243 .243 .243 1-1081 - .497 .243 .243 .243 .243 .244 .243 .244 1-1082 - .1075 .147 .245 .344 .003 .465 .243 .248 .244 .248 1-1082 - .1151 .031 .044 .243 .243 .243 .243 .243 .243 1-1082 - .1231 .031 .245 .245 .243 .244 .243 .244									-		9.362	-002
1.114	***				_			300	3.413	0 09.1	000	4.084
1 1.075 -1 376 844 647 -2 1.316 2.546 -2 3.66 3.347 3.30 -1 1.531 1 1.	1.144							.334	3.378		1.058	6-028
4.677 - 2 .477 .744 2.31 - 2 1.356 .096 2.28 - 2 .193 3.320 1.97 - 1 1.784 1.3 .669657 .652 - 3 1.371 .081 2.28 - 2 .117 3.295 1.43 - 1 1.631 1.3 .689657 .622 - 3 1.371 .081 2.23 - 2 .160 3.273 1.431 1.631 1.3 .680657 .622 - 3 1.359 .053 3.53 - 2 .264 3.164 8.94 - 2 2.565 2.7 .697 .147 4.36 - 3 1.643 .008 1.00 1.28 - 1 3.122 2.7 .187 - 3 .187 - 3 .1643 .009 1.666 - 1 1.239 2.173 1.28 - 1 3.122 2.7 .4,137 - 3 .1443 .009 1.666 - 1 1.239 2.173 1.67 - 1 3.312 2.2 .1341 .001 .941 - 5 1.442 .000 4.35 - 2 2.193 1.67 - 1 4.66 - 2 2.193 </td <td>1.073</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>.366</td> <td>3.347</td> <td></td> <td>1.531</td> <td>4.555</td>	1.073							.366	3.347		1.531	4.555
1	. 4.677								3.320		1.784	4.302
	. K-693							111.	3.295		1.951	4.135
2 4.799 - 2 1.075 .147 4.36 - 3 1.424 .028 1.30 - 1 .523 .2.789 1.28 - 1 .2.555 1.154 1.25		_						744	3.273		2.079	4.007
7 1.7.7 1.1.14 1.001 2.45 -4 1.443 1.009 1.66 -1 1.2.59 2.173 1.19 1.3.335 1.79 1.3.335 1.	7.75							. 523	2.789		2.565	3.521
3 9.440 -4 1.214 .007 5.41 -5 1.452 .000 4.35 -1 2.193 1.220 1.87 -1 4.832 .000 4.35 -2 2.193 .220 4.46 -2 5.859 .000 1.84 -4 3.517 .001 9.91 -7 1.452 .000 1.84 -4 3.517 .001	27.4	_						1.259	2.173		3.836	2.250
0 1.717 -4 1.721 .001 9.91 -7 1.452 .000 1.84 -4 3.17 .001	4.440	_						2.193	1.220		4.832	1.254
		_						3.173	022		5.859	.227

5.896 4.976 4.946 4.299 4.135 4.007 3.521 2.250 1.254 2.250	5.647 4.594 4.594 4.135 4.135 4.135 2.296 2.250 2.227	5.493 4.858 4.293 4.293 4.293 5.203 5.256	5.345 4.811 4.296 4.139 4.007 4.007 5.21 2.256 1.254 1.254 1.254
. 000 . 920 1. 353 1. 761 1. 885 2. 375 2. 375 2. 375 4. 645 4. 645 5. 669	. 000 . 742 1.123 1.952 1.512 1.640 2.663 3.397 6.393 5.645	. 000 . 635 . 982 . 982 1 . 972 1 . 972 1 . 972 1 . 973 4 . 239 5 . 286 5 . 491	
1.36 6.01 - 1 1.67 - 1 1.69 - 1 1.69 - 1 1.79 - 1 1.79 - 1 1.67 - 1 1.67 - 1 1.67 - 1	5.09 -1 1.00 -	6.73 -1 2.657 -1 1.651 -1 1.651 -1 1.181 -1 1.181 -1 1.87 -1 1.87 -1 1.87 -1	2.55 1.1 1.58 1.1 1.58 1.1 1.58 1.1 1.58 1.1 1.58 1.1 1.69 1.2 1.50 1.2 1.5
3.413 3.348 3.347 3.347 3.250 3.273 3.148 2.189 2.189 2.20	3.413 3.3413 3.347 3.347 3.220 3.233 2.189 2.189 2.20 0.220	3.413 3.347 3.347 3.347 3.295 3.295 3.248 2.173 1.220 .220	3.413 3.347 3.347 3.347 3.245 3.245 2.173 3.220 2.20
	.300 .334 .346 .193 .117 .160 .264 .253 .2193 3.193	.000 .036 .036 .117 .117 .166 .264 .264 .264 .264 .364 .364 .364	
3.66. 9.	3 2 3 3 4 5 6 6 6 7 5 7 6 7 6 7 6 7 6 7 6 7 6 7 6	3.60 2.59 2.59 2.59 2.59 2.59 2.59 2.59 2.59	7 - 2 - 2 - 2 - 2 - 2 - 2 - 2 - 2 - 2 -
12.262 5.34 5.34 1.34 1.03 1.03 1.03 1.03 1.03 1.03 1.03 1.03	1.013 442 4218 4218 4045 4041 4061 6003 6003 6003 6003 6003	. 396 . 396 . 196 . 127 . 127 . 098 . 098 . 009 . 009	.111 .349 .192 .192 .094 .003 .009
.000 .746 1.025 1.128 1.128 1.259 1.259 1.252 1.252	.600 .570 .755 .883 .918 .950 .960 .964 1.010 1.012	. 0000 . 0554 . 732 . 732 . 738 . 850 . 850 . 850 . 850	.500 .361 .518 .587 .658 .658 .658 .701 .710
1.38 1.38 1.37 1.37 1.38 2.23 5.23 5.23 5.23 5.35 5.35 5.35 5.41 5.41 5.41 5.41 5.41 5.41	9 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 -	2.08 -1 1.08 -1 1.08 -1 2.01 -2 5.26 -3 5.26 -3 5.26 -3 5.26 -3 5.26 -3 6.85 -3 6.81 -5 9.41 -5	5.25 1.10 1.00
1.222 i.044 .459 .466 .744 .744 .167 .067 .007	1.22 1.024 .959 .959 .956 .744 .147 .047 .007	1.222 1.044 .959 .959 .952 .952 .957 .067	1.222 1.084 . 959 . 846 . 846 . 852 . 320 . 147 . 067
. 0000 . 138 . 264 . 376 . 569 . 901 . 1007 1. 154 1. 154 1. 154 1. 154	.000 .136 .136 .376 .477 .569 .901 1.075 1.191 1.191 1.214		.000 .138 .263 .264 .477 .569 .901 1.075 1.154 1.191 1.221
1.444 -1 1.184 -1 1.184 -1 1.073 -1 1.073 -2 4.041 -2 1.050 -2 4.737 -3 1.212 -4	1. 112 -1 1. 112 -1 1. 113 -1 1. 113 -1 1. 113 -1 2. 2.44 -2 1. 0.50 -2 1. 0.50 -2 1. 0.50 -2 1. 0.70 -2	1.446 -1 1.1312 -1 1.1073 -1 1.1073 -2 4.881 -2 4.050 -2 4.737 -3 4.737 -3	1.444 1.018 -1 1.018 -1 1.018 -1 4.881 -7 4.881 -7 4.881 -7 4.732 -3 4.732 -3 4.732 -3
247444343444	34444545463		3-44.373644

The state of the s

١	¥ 1	Rayleigh	Rayleigh	Payleigh	able 3.4.	Param	Parameters at	0.32 Microns	rons				
Rge		atten.	optical	optical	atten	parical	Aerosol optical	Ozone	Ozone	Ozone	Ext.	Ext.	Ext.
, ,	,	coeff.	thick.	thick.	coeff	thick.	thick.	coeff.	optical thick.	optical	coeff.	optical	optical
E E	<u>.</u>	(km ⁻¹)	(q- 0)	(Å.	(ka.'1)	(u-h)	(F)	(km-1)	(u -0)	(}- ()	(kg ⁻¹)	(0-h)	(h-e)
ا ځ	£	g 4	,,1⊷	두H	_n a.	⁺ e-	- <u>-</u> d	e* 5	, r e	- M	f ext	ext	ext
	7		000.	124.	3.30	03',	2.840	10.00					j
		1- 146-6	•105	6/8.	0 00-1	1.949	006	7.44 - 3	200	505	9.6	. 000	4.080
	`	2- 020°4	351.	. 178		2.530	916	2-64-4		900		2.026	4202
	-	4.148 -7	-285	749.		2.704	9	2.26 - 4	700	567		2007	1.003
	•	7-567 -7	.362	.505	2-00-2	2.755	*60*	2.03 -3	010	293	7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	4. 128	25.6
•	٠:	7- 46.0	32	4.65		2.711	.078		710	167.		3.215	865
ı	: :	7- CD7	* 0 .	.263		5.199	.051	3.14 -3	.324	.280		3.506	574
	: ?	7-149-7	918.	1 1 0		2.822	170-	8.93 -3	.355	.248		3.693	.337
	: :	1000	2 2	140.		2-640	*00*	1.47 -2	.110	.193		3.827	.253
	: =	7.586 -4	625	• 70.0		958-7	•003	1-62 -2	-195	901.		3.945	.135
		4- 207	725	5 6		648-7	000	3.87 -3	.284	070		4.055	•025
	!		•	5		648.7	000	1-67 -4	.103	000•		4.079	100
	•		000		20.00	0.00		;					
		1- 104.4	607		7.73	200	2.50	3-20 -3	200	.303		900	3.291
	`	9.070 -/	7	50	2.16 -1	1 2 3 5	2175	2.43 -3	.003	.300	4-25 -1	1.446	1.835
	-	8-148 ->	-285	749	7.58 -2	21.7	1 20		900.	. 298		1.977	1.304
•	•	1-147 -1	. 362	.505	2-42 -2	7.6.	650		907.	567		2.205	1.076
~	~	6-544 -2	.437	.445	7.94 -3	1.972	670			567		2.330	156.
	3	3.705 -2	*P9*	.243	5.02 -3	2.000	0.51		324	767		114.7	698.
	ζ.	1-745 -7	. 816	111	4-19 -3	2.023	.027		1355	268		20.7	
	7 .	£- /05.	. H76	.051	2-36 -5	2.041	÷00•		0110	.193		3.028	
	; ;	2. 14. 1.	*05	*20.0	4- 15-9	2.048	-003	1.62 -2	.195	.108	2-06-2	3.146	.135
	2,	4- 500-6	. 63		9 - 1 - 0	0000	000.		- 284	.020		3.256	.025
			:			0000	000-		.303	000		3.280	100
	7	1- 850-1	700	766.		Q			;	,			
	-	7- 144.6	.105	.623		420-1	7097	3.64	900	.303	1.78 O	000	2.863
	`		.199	.728		1.377	.255	2.63 -3	000	300		1.133	1.730
	•	8.1 td	507-	.649		1.458	.134	2.24 -3	306	295		1.703	1.260
7		7- 7-1	<u>ک</u> و:	\$05.	2-31 -2	1.540	•093	2.03 -3	010.	.293	9. de -2	1.917	1.0
۲	` :)	,		1.554	.076	1.36 -3	716.	.291		1.998	. 865
	<u>.</u>	1.745	418			1.522	150-	3-14 -3	.324	.280		5.289	.574
	?		47.	.051		1.623	500	8-43-3	• 155	.248		2.476	.387
	;		₹ 0¢•	•70-		1.629	600	2- 14-1	901	661.		2.610	•253
	٠.		- 922	500°		1.632	000	3.87 - 5	787	020		2 939	
			1/4-	100.		1.632	.000	1-57 -4	.303	000	2.60 -4	2.862	100
			;	!									
	o -	- 840.1	000.	.927	1.33 O	003.	1.370	3.23 -3	.300	.303	1.44	000	2,601
	. ~		501.	. 863		. 633	.537	2.13 -3	.303	•300		3	1.660
			A 2 0	2		1-132	-238	2.63 -3	900.	.298		1.337	1.264
	•	7- 5-6-7	(4)	40.4		1-239	181.		.304	-585		1.533	1.068
L.	•		, ;	5,7		1.216	760.	2.03 -3	010	-293		1.651	.950
•	7		\$20.	.243		1.370	8.00		2315	.291		1.736	-865
	2		. dlb	111.		1.543	770	C + 77 - X	*36.	3,50		2.027	.574
	? :	7.407 -4	016.	140.	7-36 -3	1.361	600	1-47 -2	77		2- 51 - 2	2.214	.387
	; ;		Ď.	•024		1.367	£00°	1.62 -2	.195	-108		2.466	135
	: 7	41 677	775	, o	5-21 -5	1.370	000	3.67 - 3	.284	.020		2.576	.025
			:	•	- 66.4	1.570	000	1-67 -4	.303	000.	5.60 -4	5.600	.001

e nen meg Appelations

2.422 1.611 1.051 1.066 .950 .957 .387 .259 .135	2.187 1.543 1.543 1.062 1.062 2.649 2.649 2.53 2.135 2.000	2.042 1.500 1.222 1.059 .865 .865 .387 .387 .387	1.903 1.450 1.209 1.054 1.054 1.054 1.054 1.055 1.055 1.055
. 000 . 11. 17. 11. 17. 11. 17. 11. 17. 11. 17. 11. 17. 11. 17. 11. 17. 11. 17. 11. 17. 17		. 000 . 943 . 923 . 923 . 923 . 923 . 1 . 656 . 1 . 656 . 1 . 656 . 1 . 900 . 2 . 017	
1.22 5.15 - 1 2.46 - 1 9.68 - 2 7.59 - 2 7.59 - 2 3.68 - 2 2.04 - 2 2.04 - 2 5.06 - 3	5 5 5 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	6 6 9 6 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7
. 299 . 299 . 299 . 299 . 299 . 299 . 199 . 199 . 199 . 199	. 200 . 200	.303 .303 .203 .203 .203 .203 .203 .203	. 303 . 309 . 298 . 293 . 293 . 294 . 294 . 193 . 193 . 108
00000000000000000000000000000000000000			200 200 200 200 200 200 200 200 200 200
3.20 2.50 2.50 2.50 2.60 2.60 3.10 3.10 3.10 3.10 3.10 3.10 3.10 3.1	2.42 2.43 2.43 2.43 2.43 1.98 1.98 1.94 1.67 1.67 1.67 1.67 1.67	2.20 - 2 2.693 - 3 2.693 - 3 2.264 - 3 2.264 - 3 2.264 - 3 3.167 - 3 3.167 - 3 3.167 - 3 3.167 - 3 3.167 - 3	2.000 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
1-191 -488 -226 -126 -078 -078 -007 -009 -009		. 122 . 136 . 136 . 122 . 031 . 031 . 003 . 003	
			5533 5533 5533 5533 5533 5533 5545 5545
11.1 14.12.1 14.12.1 14.12.1 14.12.1 14.13.1 1	8.23 -1 3.25 -1 5.08 -2 7.09 -2 7.09 -3 6.34 -3 6.34 -3 6.51 -5 6.51 -5	7 - 7 - 7 - 7 - 7 - 7 - 7 - 7 - 7 - 7 -	5.00 6.18 6.18 6.18 6.18 6.18 6.18 6.18 6.18
. 927 . 728 . 728 . 545 . 545 . 645 . 651 . 651 . 605	. 927 . 823 . 728 . 642 . 645 . 645	. 622 . 623 . 728 . 642 . 642 . 643 . 643 . 643 . 643 . 643 . 643 . 643 . 643 . 643	. 923 . 642 . 642 . 642 . 643 . 644 . 644
. 105 . 105 . 105 . 285 . 385 . 415 . 816 . 906	. 100 . 100	.000 .100 .199 .385 .385 .385 .385 .906 .906	
1.098 -1 9.962 -2 9.962 -2 8.168 -2 6.597 -2 7.967 -3 7.967 -3 7.597 -3 7.597 -3	1,098 -1 6,082 -2 6,182 -2 7,342 -2 3,705 -2 1,705 -2 1,705 -2 1,705 -3 7,594 -3	1,000 9,900 9,900 1,000	1,008 2,922 2,922 2,922 2,923 2,924
(1)	**************************************	2	3

				1	Table 3.5.	. Parameters	eters at	0.34 Mic	crons				
į	Alt.	Rayleigh	Rayleigh	Ray leigh	Aerosol	Aerosol	Aerosol	Ozone	Ozone	Ozone	Ext.	Ext.	F. 4
Ļ		etten.	optical	optical	atten.	optical	optical	absorp.	optical	optical	coeff.	ontical	
9	3		thick.	thick.	Coeff.	thick.	thick.	Coeff.	thick.	thick.	1	thick.	thick.
Ì		ļ		r	1	(11-0)	(<u>•</u> -u)		(u-n)	((.	(0-h)	(•-e)
>	Æ	es ⁽⁴	_r ••	1.th	∞ ₽.	, a ,	- <u>c</u>	в3	ř.	. K	beat	, ext	ext
	,			;									
	> -	7- 707 -2	3		9.14	99.	2.675	2.28 -4 2.56	000	-022	3.22	00	3.414
	~	6-978 -2	*	X	2.62 -1	2,317	101	1.88 -4		120		2 526	1.509
	-	6-303 -2	.221	.497	1-41	2.536	.139	1-60 -4	.001	.621		7.757	
•	•	2-099%	 	.437	2- 45-2	2.565	060	1.45 -4	.001	.021		2.866	848
7	•	5-105	**	. 343	7.63 -3	2.600	.075	1-41 -4	100-	.021		2.935	624
	-	2-1997	625	=	4.8 2 -3	2.627	640-	5-54 -4	-305	.020		3.157	.257
	23	2-066-1	169	980	4-03 -3	2-649	•026	6.36 -4	•00•	910.		3.284	.130
	? :	2. 27		100	2.26 -3	2.666	600	1.05 -3	800.	+10-		3.352	.062
	; ;				00.0	719.7	500.	6- 61-1	*1C*	800		3.385	-029
	ò	7-119 -5	711.	700	9-14-7	2.675	88	1-19 -5	.322	000	8-40	3.413	900.
	-	403	ě	;	;	;		;					
	• -	7- 101-1	8		20.7	000	1.926	2.28 -4	000	-052		8	2.665
	~	6-978 -2	*	563		1.643	. 26.5			12		1.334	1.331
	•	6-303 -2	.221	2 R. 9		1.754	1 33			120		910	200
1	•	2-080-5	.280	.437		1-837			100	120		2.013	000
~	•	2-101-5	.334	-383		1.851	.075		100	170.		2.186	470
•	2	2-140-5	675	=		1.678	640		200	.020		2.409	25.7
	<u>.</u> :	1-350	1691	100		1-900	•056		.304	910		2.535	-130
	2 :	6-104 -5	-678	. 034		1.918	60,		.008	•10 ·		2.603	-062
	; ;		7.1		91 00 9	1.924	.003		•114	900-		2.637	-029
	3	2- 611-2	.117	8	9.14 -7	1.926	3 9	1-14-5	220	1000	4. 13. 4. 4. 6. 15.	7.000	8 8
													5
	•	6.497 -2	000	.711	0 %:1	000	1.534	2-28 -4	,300	.022		8	
	-	7- 101 -2	190	**		96.	.574	2-09-4	000	-021		3	1.232
	~ •	6-978 -2	<u> </u>		7 2 :	1.291	.243	1-88 -4	.300	.021	2.56 -1	1-4-6	.827
	. ,	5-640 -2	780	7		904-1	821-		100.	-021		1.627	.646
7	•	2 SOL 2	7.			1.459	20.0		100	120		1.726	.547
•	2	2- 190-2	83.			1.415	640		2003	050		2.016	25.7
	<u>.</u>	7- 956-1	.631	*		1.568	-026		+000	.018		2.143	.130
	2 (- 22	8			1.525	600		900	*01		2.211	-062
	÷	7 E87	.713	400		1.534	000	2-76 -4	320	500		2.24	-029
	3	7.119 -5	.117	ã.		1.54	000		.322	000	8-40 -5	2.272	100
	3	4-407-7	000	7117	1.24 0	000	1.084	4				;	!
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					Table 3.6		Parameters at	0.36 Microns	rons				
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	. ~		121-			2.16	2.20		9	ē.		2.346	62.
	•		-220	1		7.431	981			3		66.2	. 26.4
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					Table 3.9		Parameters a	t 0.45 Microns	crons				
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	•		.069	.155		1.674	977		960	8		000	.403
c	٠		-087	-136		1.911	700		000	3 8			200
4	•		5	617-		1.922	-056		2000	8		8	17.
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	3		1410	7.7		1.959	•020		.000	100-		2.156	3
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	42		-225	100		1.978	200		100	000-		2-195	900
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					Table 3.1	l2. Para	Parameters	at 0.60 M	Microns				
Met. A	<u>:</u>	Rayleigh	Rayleigh	Rayleigh	Aerosol	Aerosol	Aerosol	Ozone	Ozone	Ozone	Ext.	Ext.	Ext.
•		coeff	thick	thick.	coeff.	thick.	thick.	coeff.	thick.	thick.	. Lett.	thick.	optical thick.
- 	ê	(km-1)	(4 -0)	Ė	(km ⁻¹)	(q-0)	÷	(km ⁻¹)	(u-0)	(<u>+</u> -	(km)	(q-0)	() ()
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•	•		.015	• 0.		1.328	-142		100.	**0		1.344	.280
			- 021	# to .		1.425	•040		100.	.043		1.447	.177
•	•	4- 444.4 4- 444.4	200	750		1.454	950.		200.	.043		1.483	141
7			160	810.		1.440	0.0		303	100		1.534	060
	•		190.	HO0.		1-454	410.		80C-	.036		1.563	.061
•	,		- 062	*00*		1.505	500-		. 110	.028		1.586	.039
•			9	50c.		505-1	205.		.329	910.		1.604	610.
- Å	• •	7.017 -5 h.017 -6	900.	000	2- 21-4 2- 51-4	1.510	000	2.43 -5	345	000	3.20 -5	1.624	.000
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•		6-101-3	800 ·	190	1-28-1	970	265-	3.47	200	440		246	757
-		6- 15t -4		6		200	680		100			5	173
•			.027	645		1.035	440.		305			1.063	141
~	•	4-404-4	.032	110.		1.043	190-		700-	.043		1.077	.127
-	•		140.	.018		1.060	.030		.303	140-		1.114	040.
- - '		•	٠.	800.		*CO-1	-016		80C.	.036		1.143	190.
` `	٠,	3- 414 ·	\$60.	*00		1.085	5005		910	970-		1.165	.034
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	٠.	B-156 -3	900	690.		000-	.470		000.	• 042		000	* 96*
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			170.	840.		16/	610.		101	.063		813	170
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`	~		.065	•00•		. 465	5000		910.	.028		946	.034
•	•		190.	200.		999.	700-		.324	-016		* 364	610.
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	7	4-130 -4	000-	.00		000-	.132		.000	5 00 -	6.8 -1	900	.846
•	4	1.401 -5	BOO.	190.		.434	96₹*		.300	**0.		. 443	.403
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13. ABSTRACT					
An examination of the surface meteorological ra (2) the lower and upper vi logical ranges 1.2 km and ranges are selected for de An aerosol scale height is puted aerosol attenuation viously published attenuatialitude.	nge generally affi sibility limits of 15 km respective veloping uv, visi derived for each	ect a mixing lay a haze regime bely. Within these ble, and ir aero meteorological presented as tab	er to 5 are des elimits sol atto range ulation	i km fined eiglenuat Fi s wh	by meteoro- nt meteorological cion coefficients. nally, the com- ich include pre-
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Haze transmission in UV Haze transmission in Visible Haze transmission in IR	 NK B	ROLE	WT
Haze transmission in Visible			
Vertical aerosol attenuation through haze Haze attenuation model Light scattering in atmospheric haze			

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